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**FEASIBILITY OF USING THE MASW METHOD TO DEFINE A
SINKHOLE IMPACT AREA AT
CALVERT CLIFFS NUCLEAR POWER PLANT, MARYLAND**

by

Jianghai Xia
Richard D. Miller
Choon B. Park

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Kansas Geological Survey
1930 Constant Avenue
University of Kansas
Lawrence, KS 66047-3726

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Jianghai Xia
Richard D. Miller
Choon B. Park

Kansas Geological Survey
1930 Constant Avenue
Lawrence, Kansas 66047



Final Report to

Earth Resources Technology, Inc.
Jessup, Maryland

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Jianghai Xia
Richard D. Miller
and
Choon B. Park

of the

Kansas Geological Survey
1930 Constant Avenue
Lawrence, KS 66047

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Earth Resources Technology, Inc.
8305 Stayton Drive, Suite B
Jessup, MD 20794

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Abstract

A sinkhole developed at Calvert Cliffs Nuclear Power Plant, Maryland, early this year. Approximately 700 tons of soil were used to fill in the sinkhole to prevent further damage. A high-frequency surface-wave survey was conducted in hopes of defining the sinkhole impact area. Acquisition of surface-wave data was in an extremely noisy environment. The S-wave velocity field calculated from surface-wave data showed chimney-shaped low-velocity anomalies that were directly related to the sinkhole. Based on S-wave velocity field maps, a potential sinkhole impact area was tentatively defined. S-wave velocity field maps also revealed, depending on the acquisition geometry, one side of the water tunnel of the power plant.

Introduction

Elastic properties of near-surface materials and their effects on seismic wave propagation are of fundamental interest in groundwater, engineering, and environmental studies. S-wave velocity is one of the key parameters in construction engineering. As an example, Imai and Tonouchi (1982) studied P- and S-wave velocities in an embankment, and also in alluvial, diluvial, and tertiary layers, showing that S-wave velocities in such deposits correspond to the N-value (Craig, 1992), an index value of formation hardness in soil mechanics and foundation engineering.

Surface waves are guided and dispersive. Rayleigh (1885) waves are surface waves that travel along a “free” surface, such as the earth-air interface. Rayleigh waves are the result of interfering P and S_v waves. The particle motion of Rayleigh waves moving from left to right is elliptical in a counter-clockwise (retrograde) direction. The motion is constrained to the vertical plane consistent with the direction of wave propagation (Babuska and Cara, 1991, p. 30). Longer wavelengths penetrate deeper than shorter wavelengths for a given mode, in general exhibit greater phase velocities, and are more sensitive to the elastic properties of the deeper layers (Babuska and Cara, 1991, p. 30). Shorter wavelengths are sensitive to the physical properties of surficial layers. For this reason, a particular mode of surface waves will possess a unique phase velocity for each unique wavelength leading to the dispersion of the seismic signal.

The S-wave velocity can be derived from inverting the phase velocity of the surface (Rayleigh and/or Love) wave (Dorman and Ewing, 1962). For the case of a solid homogenous half-space, the Rayleigh wave travels with a velocity of approximately $0.9194v_s$, where v_s is the S-wave velocity in the half space (Sheriff and Geldart, 1982, p. 49) and is not dispersive. In the case of one layer on top of a solid homogenous half-space, dispersion of the Rayleigh wave occurs when the wavelengths of the Rayleigh wave are in the range of 1 to 30 times the layer thickness (Stokoe et al., 1994). Stokoe et al. also show that the Rayleigh wave travels with a velocity of approximately $0.9194v_l$ (v_l is the S-wave velocity in the layer) when the wavelengths of the Rayleigh wave are shorter than the layer thickness. For the wavelengths longer than 30

times the layer thickness, the Rayleigh wave phase velocity is approximately equal to $0.9194v_2$ (v_2 is the S-wave velocity in the half space).

Ground roll is a particular type of Rayleigh wave that travels along or near the ground surface and is usually characterized by relatively low velocity, low frequency, and high amplitude (Sheriff, 1991, p. 143). Stokoe and Nazarian (1983) and Nazarian et al. (1983) presented a surface-wave method, Spectral Analysis of Surface Waves (SASW), that analyzes the dispersion curve of ground roll to produce near-surface S-wave velocity profiles. SASW has been widely applied to many engineering projects (e.g., Sanchez-Salinero et al., 1987; Sheu et al., 1988; Stokoe et al., 1989; Gucunski and Woods, 1991; Hiltunen, 1991; Stokoe et al., 1994).

Inversion of dispersion curves to estimate S-wave velocities deep within the Earth was first attempted by Dorman and Ewing (1962). Song et al. (1989) related sensitivity of model parameters to several key parameters by modeling and presentation of two real examples where surface waves were used to obtain S-wave velocities. Turner (1990) examined the feasibility of inverting surface waves (Rayleigh and Love) to estimate S-wave and P-wave velocities. Dispersion curves are inverted using least-squares techniques in SASW methods (Stokoe and Nazarian, 1983; Nazarian et al., 1983). Rix and Leipski (1991) examined the influence of the number of dispersion data, the maximum wavelength, and distribution of dispersion data with wavelength on the accuracy and resolution of S-wave velocity profiles.

The Kansas Geological Survey conducted a research project to estimate S-wave velocity from ground roll. The project consists of three phases: acquisition of wide band ground roll, creation of efficient and accurate algorithms organized in a basic data processing sequence designed to extract Rayleigh wave dispersion curves from ground roll, and development of stable and efficient inversion algorithms to obtain S-wave velocity profiles.

Based on our experience, the fundamental-mode phase velocities, when calculated with high accuracy (e.g., Xia et al., 2000b), can generally provide reliable S-wave velocities ($\pm 15\%$). The accuracy can be increased if higher modes of surface waves are available and included in the inversion process (Xia et al., 2000a). Recent advances in the use of surface waves for near-surface imaging incorporated Multi-channel Analysis of Surface Waves (MASW) (Park et al., 1996a and 1996b; Xia et al., 1997; Park et al., 1999b; Xia et al., 1999; Xia et al., 2000b) with a standard CDP (common depth point) roll-along acquisition format (Mayne, 1962) similar to conventional petroleum exploration data acquisition. Combining these two uniquely different approaches to acoustic imaging of the subsurface allows high confidence, non-invasive delineation of horizontal and vertical variations in near-surface material properties (Miller et al., 1999, Xia et al., 2000c).

A sinkhole (see cover photo) developed at Calvert Cliffs Nuclear Power Plant, Maryland (Figure 1). The subsurface drainage system is a possible cause of the sinkhole. The original groundwater level was 15 to 20 ft above the sea level in the plant area; however, a permanent pipe drainage system surrounding the plant maintains the groundwater below 10 ft above sea level. At the same time that the drainage system lowers the groundwater level, soil may wash away so that over time void(s) may develop.

Southwestern Maryland

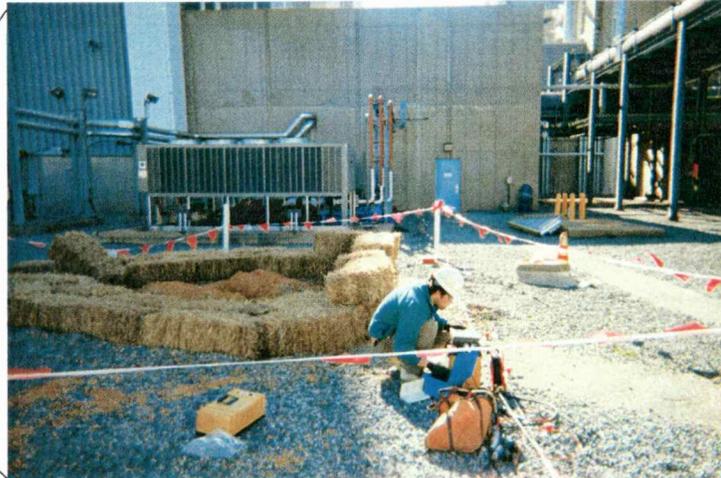


Figure 1. A sinkhole surrounded by straw bales at Calvert Cliffs Nuclear Power Plant, Calvert County, Maryland. The structure at the back of the photo is the turbine building. The transformer is located to the right in the photo.

The area potentially impacted by the sinkhole needs to be defined to prevent further damage and assure the safety of the plant. Voids can be difficult to detect from surface investigations and many times provide an expensive challenge to detection with hit and miss drilling. Changes in shear wave velocity may be expected at the boundary between unconsolidated materials and consolidated bedrock as well as unconsolidated and fill materials. The surface-wave method was proposed to define the sinkhole impact area. This study will focus on the following goals and objectives: 1) processing noisy surface-wave data, 2) feasibility of defining a sinkhole impact area at least 50 ft deep by using MASW methods with fundamental mode data acquired with the standard CDP roll-along acquisition format, and 3) evaluation of effective depth of imaging.

The Method

Surface-wave imaging has shown great promise detecting shallow tunnels (Xia et al., 1998) and bedrock surfaces (Miller et al., 1999; Xia et al., 2000c). Extending this imaging technology to include lateral variations in lithology has required a unique approach incorporating MASW and CDP methods. Integrating these techniques provides a 2-D continuous shear wave velocity profile of the subsurface. Signal enhancement resulting from determination of a dispersion curve using upwards of 60 closely spaced receiving channels and the calculation of a dispersion curve every 2 ft or so along the ground surface provides a unique, relatively continuous view of the shallow subsurface. This highly redundant method enhances the accuracy of the calculated shear wave velocity and minimizes the likelihood the irregularities associated with an occasional erratic dispersion curve will corrupt the data analysis.

The MASW method is based on the fact that the S-wave velocity is the dominant influence on Rayleigh wave for a layered earth model, which assures us that inverting phase velocities will give us an S-wave velocity profile (1-D S-wave velocity function, V_s vs. depth). Because data are acquired in the standard CDP format, phase velocities of ground roll can be

extracted from each shot gather so that numerous 1-D S-wave profiles along a survey line can be generated. A two-dimensional vertical section of S-wave velocity can be generated by any contour drawing software.

A number of multi-channel records should first be collected in the standard CDP roll-along acquisition format. If a swept source is used, the recorded data should be converted into correlated format, which is equivalent to data with an impulsive source. Surface impact sources and receivers with a low response frequency, normally less than 8 Hz, should be chosen to acquire surface-wave data. Data acquisition parameters, such as source-receiver offset, receiver spacing, etc., should be set to enhance ground-roll signals (Park et al., 1999a).

Once the data collection is completed, phase velocities (dispersion curves) of the ground roll of each shot gather should be calculated. The frequency range and phase velocity range of the ground roll need to be determined by analyzing data along the entire line. These two ranges are very important constraints to correctly extract the dispersion curve from each shot gather. They not only help eliminate noise such as body wave, higher mode of Rayleigh waves, etc., during calculation of phase velocities (Park et al., 1999a), but also assist in defining the thickness of the layer model.

Inversion should be performed on each dispersion curve to generate an S-wave velocity vs. depth profile (Xia et al., 1999b). The inverted S-wave velocity profile should be located in the middle of the receiver spread (Miller and Xia, 1999). Initial models are a key factor to assure convergence of the inversion process. After processing more than five thousand shots of real data and countless sets of modeling data, initial models defined by Xia et al.'s algorithm (1999b) are generally converged to models that are acceptable in geology and also fit the dispersion curve in a given error range. The number of layers is chosen between five to twenty in most of our experience, based on the accuracy of surface-wave data, an investigation depth, and required resolution of S-wave velocity. The thickness of each layer varies based on the depths of interest. For example, for a geological problem with a depth of interest of 50 ft we may choose a ten-layer model with the first three layers being 3 ft thick, the next three layers 6 ft thick, the last three layers 10 ft thick, and a half space. The inverted S-wave velocity profile for each shot gather is the result of horizontally averaging across the length of the geophone spread.

Gridding algorithms, such as kriging, minimum curvature, etc., may be used to generate a two-dimensional contour map of the S-wave velocity of a vertical section. With density information, a stiffness section can be generated simultaneously. Two-dimensional data processing techniques, such as regression analysis, could be easily applied to a vertical section of S-wave velocity to enhance local anomalies.

On a 2-D S-wave velocity map, the bedrock surface is usually associated with high S-wave velocity gradients, while fracture zones, voids, buried landfill edges, and the like may be indicated by S-wave anomalies such as low velocity zones. The general procedure for generating a 2-D S-wave velocity map is summarized as follows (Figure 2).

1. Collect a number of multi-channel records in a standard CDP roll-along acquisition format. The field data acquisition parameters are chosen to enhance broad band ground roll. If a

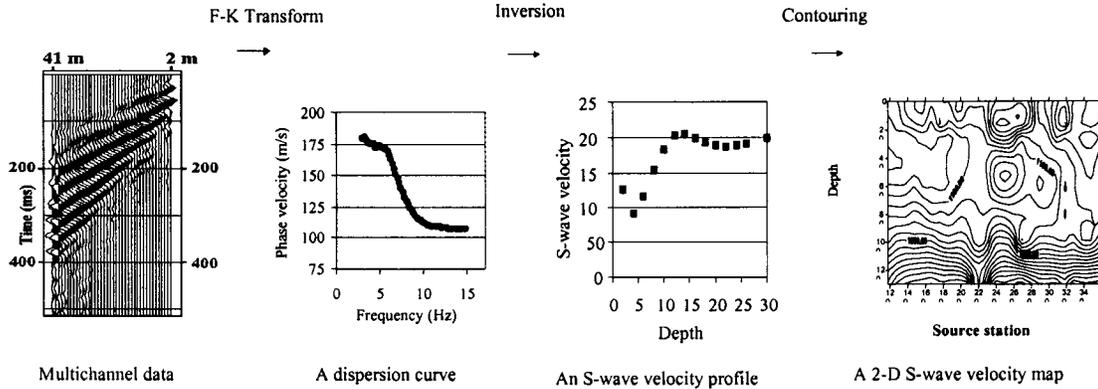


Figure 2. A diagram of construction of 2-D S-wave velocity map by the MASW method. Multichannel raw field data are acquired in the CDP roll-along format. Rayleigh wave phase velocities are extracted from the field data in the F-K domain. The phase velocities are inverted for a shear-wave velocity profile (V_s vs. depth). Finally, a 2-D S-wave velocity map is generated with contouring software.

swept source is used, the recorded data should be converted into correlated format, which is equivalent to data with an impulsive source.

2. Once the data collection is completed, calculate the phase velocity of the ground roll of each shot gather.
3. Perform inversion on the phase velocity to generate an S-wave velocity vs. depth profile at each source location.
4. Use a gridding algorithm, such as kriging, minimum curvature, etc., to generate a two-dimensional contour map of the S-wave velocity of a vertical section. With density information, a stiffness section can be generated simultaneously. Two-dimensional data processing techniques, such as regression analysis, could be easily applied to a vertical section of S-wave velocity to enhance local anomalies of S-wave velocity.

Data Acquisition

Because the depth to the drainage tunnel is 45 ft, investigation depth for this project was approximately 50 ft. Three data acquisition parameters were determined by using the investigation depth: nearest source-geophone offset, geophone interval, and length of the geophone spread (the distance from the first geophone to the last geophone).

The nearest source-geophone offset is normally chosen as half the investigation depth (Park et al., 1999b), or in this case, 24 ft. With a 24-channel recording system, the geophone spread is critical to maintain the proper resolution in the frequency-velocity domain, thereby allowing the software to accurately pick phase velocities of surface waves. The half wavelength is the effective penetrating depth of surface waves. In this project, we need to reconstruct a 100-ft wavelength component of recorded surface waves. This condition determined the geophone spread would be 92 ft with a geophone interval of 4 ft.

Earth Resource Technology, Inc. acquired the surface-wave data. Data were collected using a 24-channel StrataView seismograph with a 12-lb hammer and a 1' × 1' aluminum plate as

the seismic source (Figure 3) and a standard CDP roll-along acquisition format to generate S-wave velocity sections. Twenty-four 4.5 Hz vertical-component geophones were deployed on the same geophone interval (4 ft) with 24 ft from the source to the nearest geophone along five lines (Figure 4). Six impacts at each shot station were vertically stacked to increase the signal-to-noise ratio.

Five lines were designed around the sinkhole based on the accessibility (Figure 4) and 137 shots of surface-wave data were collected along those lines. Total line length of the surface-wave survey was 1102 ft. Total imaging length of the survey was 522 ft. The data were acquired in an extremely high noise environment. The turbine and transformer of the power plant were running during data acquisition. The noise level was as high as 100 millivolts on the noise monitor of the StrataView seismograph, which is almost 50–100 times higher than the noise level for a normal shallow seismic survey. At this noise level, no one would attempt to acquire body-wave seismic data. Surface materials consisted of gravel and asphalt. It was difficult to generate seismic energy on the gravel surface.

Data Processing

Data were processed by WinSeis[®] and SurfSeis[®], programs developed at the Kansas Geological Survey. A pre-processing flow of the surface-wave data were: a bandpass filtering (a trapezoid with four frequencies of 3, 6, 40, and 60 Hz), an automatic gain control scaling (AGC) with a time window 100 millisecond (ms), and an f-k filtering (a filtered pie defined by two lines with slopes of -3 ms/trace and -9 ms/trace, respectively). The bandpass filter and AGC were used to filter unwanted frequency components and to balance the surface-wave spectrum, respectively. The f-k filter was employed to suppress noise from the turbine and transformer (Figures 5a-5b).

SurfSeis managed to pick up dispersion curves from 13 Hz to 27 Hz for all shots (Figure 5c). Wavelengths of phase velocities ranged from 25 ft to 240 ft. A five-layer model, with thickness of 10 ft, 10 ft, 15 ft, 15 ft, and the half-space, was chosen during inversion processing. The layer thickness of the model should be controlled by the shortest wavelength of the surface waves. This low resolution model was chosen to stabilize the inversion processing because of extremely noisy data.



Figure 3. Surface wave data were collected using a 24-channel seismograph with a 12 lb hammer and a 1' x 1' aluminum plate as a seismic source.

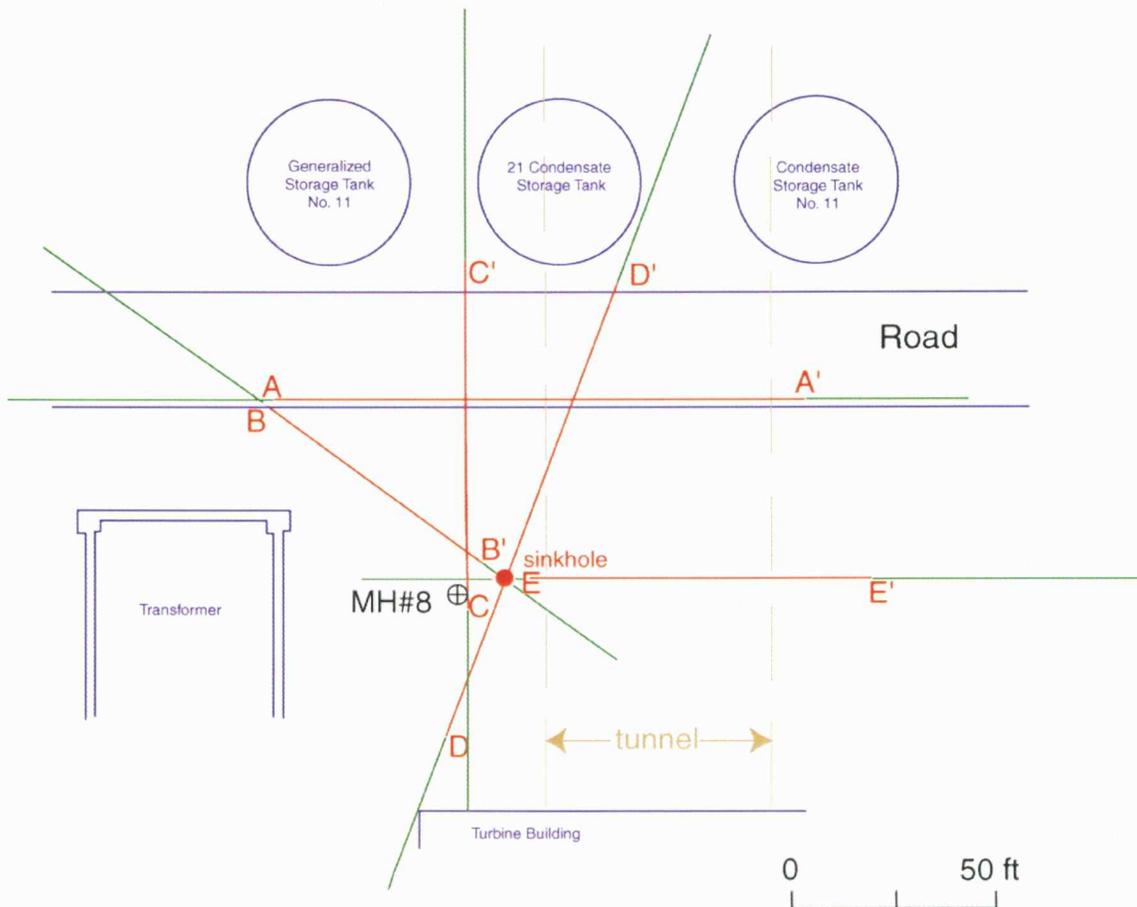


Figure 4. Location of surface-wave seismic lines. The red indicates that part of the line that was imaged by the survey.

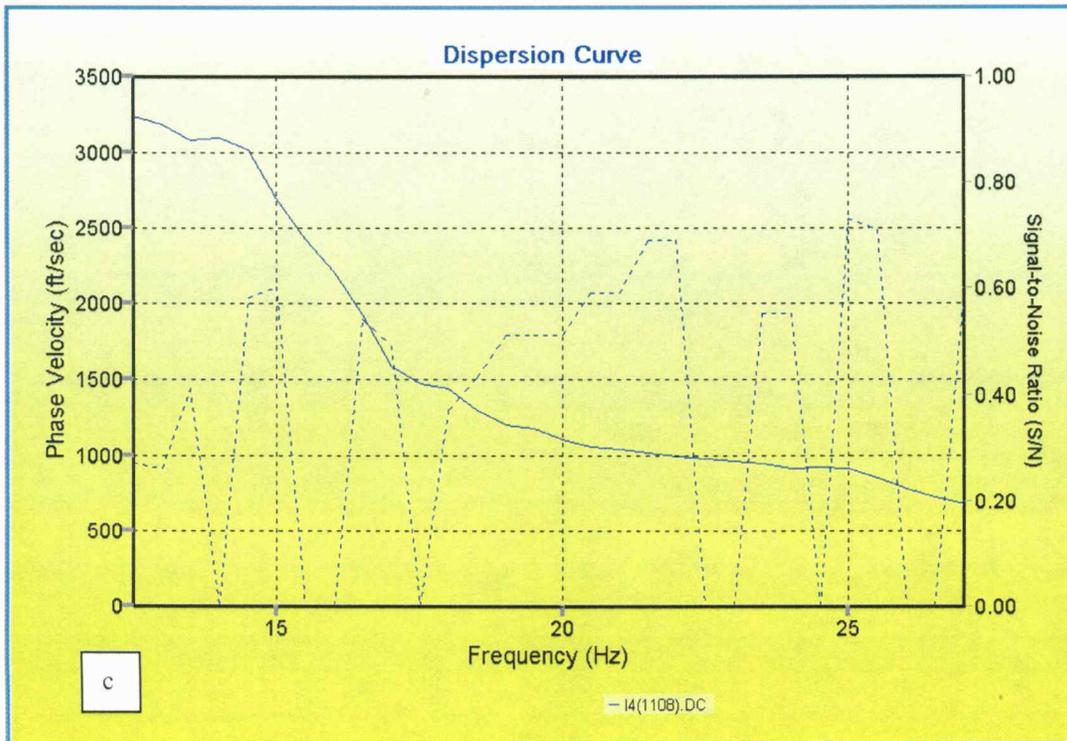
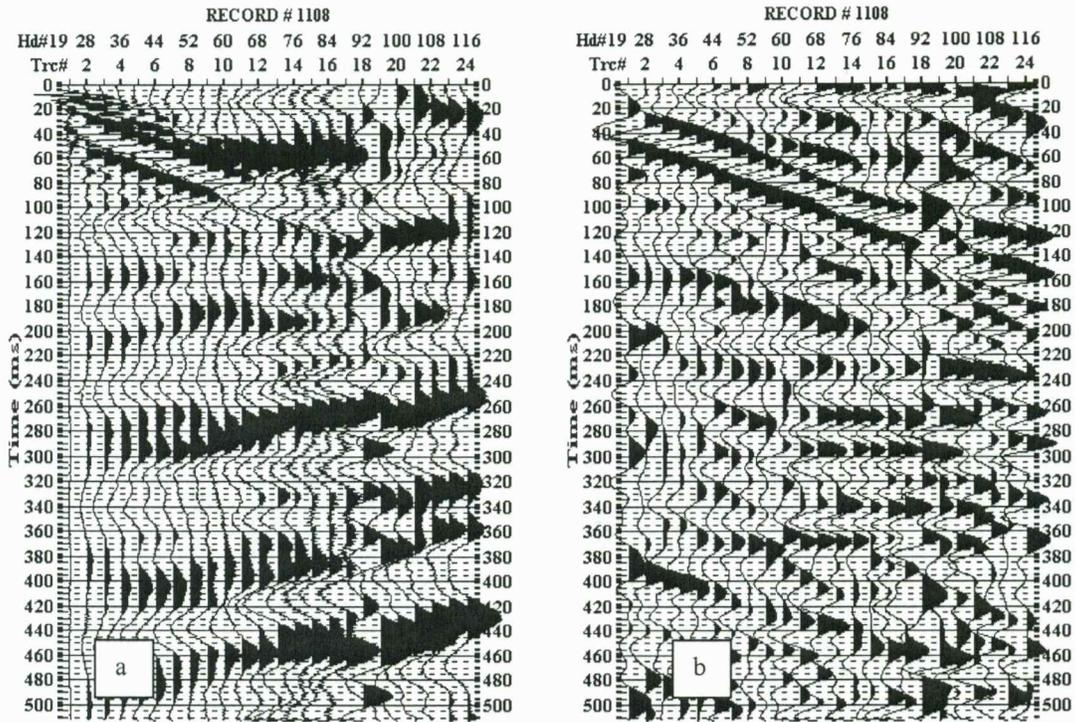


Figure 5. An example of field shot gather: a) raw data, b) data after a bandpass filtering, an AGC scaling, and an f-k filtering, and c) phase velocities (a dispersion curve, the solid line).

Phase velocities of surface waves were inverted into S-wave velocities of the layered earth model. The contouring software Surfer[®] was used to generate the S-wave velocity field (Figures 6-9).

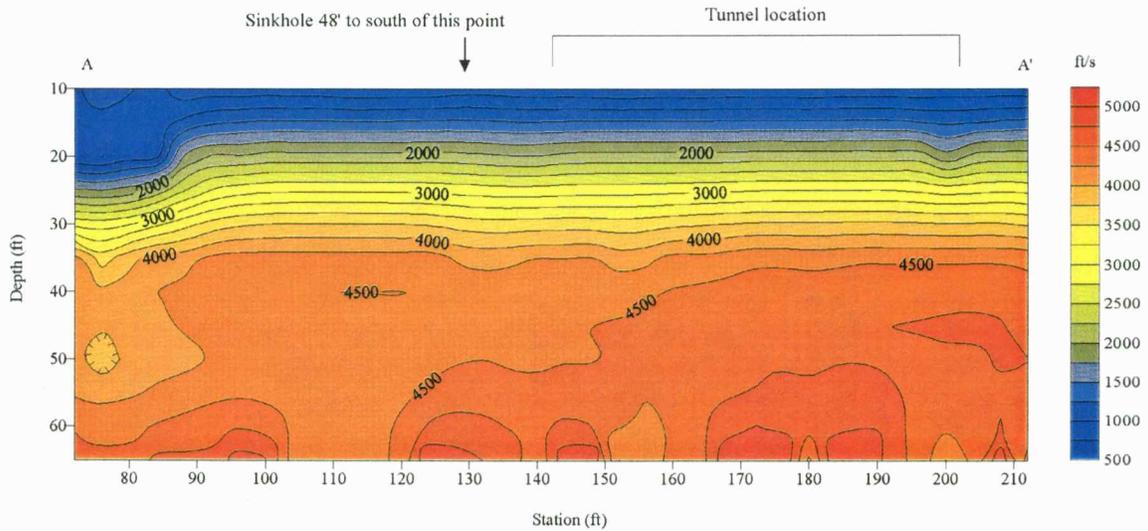


Figure 6. S-wave velocity field of line A-A'.

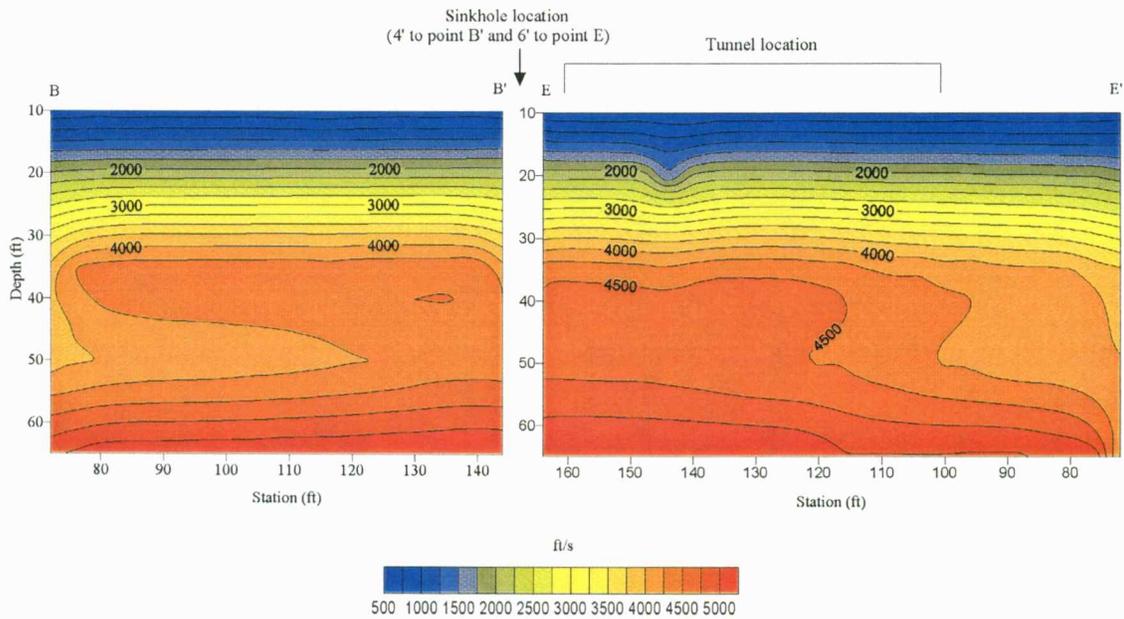


Figure 7. S-wave velocity field of lines B-B' and E-E'.

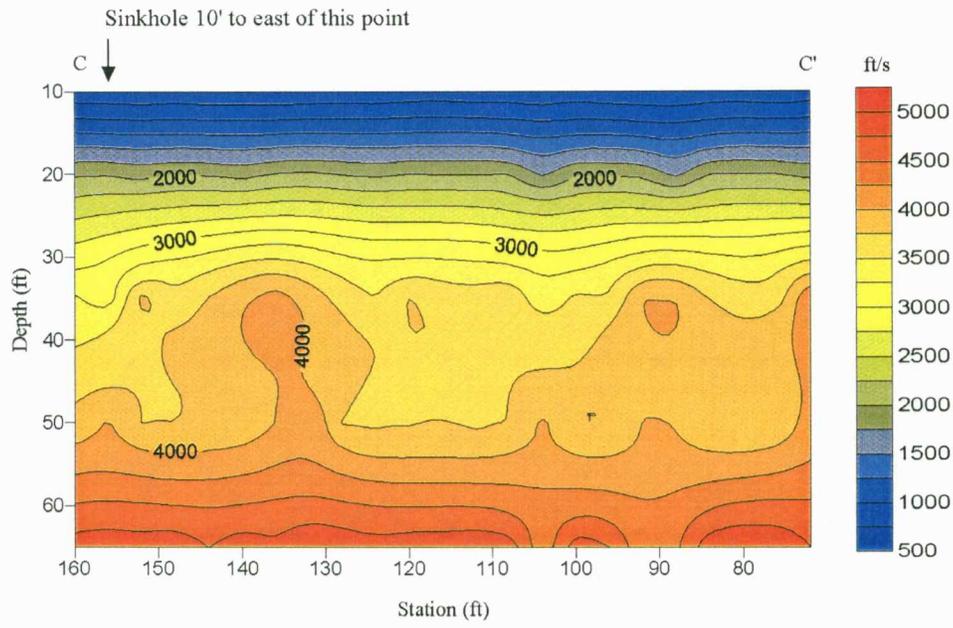


Figure 8. S-wave velocity field of line C-C'.

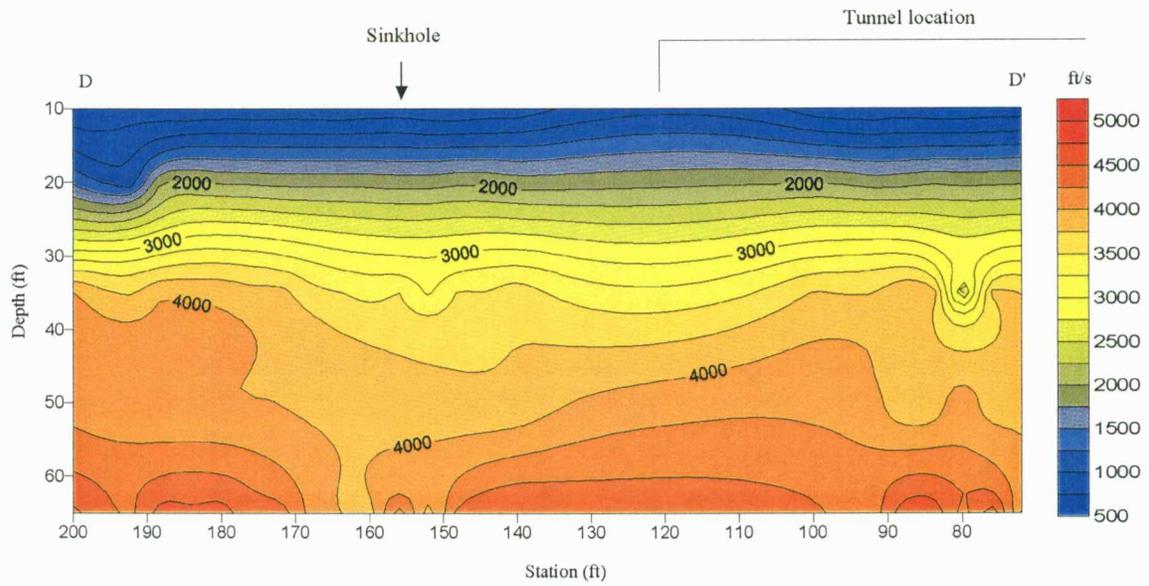


Figure 9. S-wave velocity field of line D-D'.

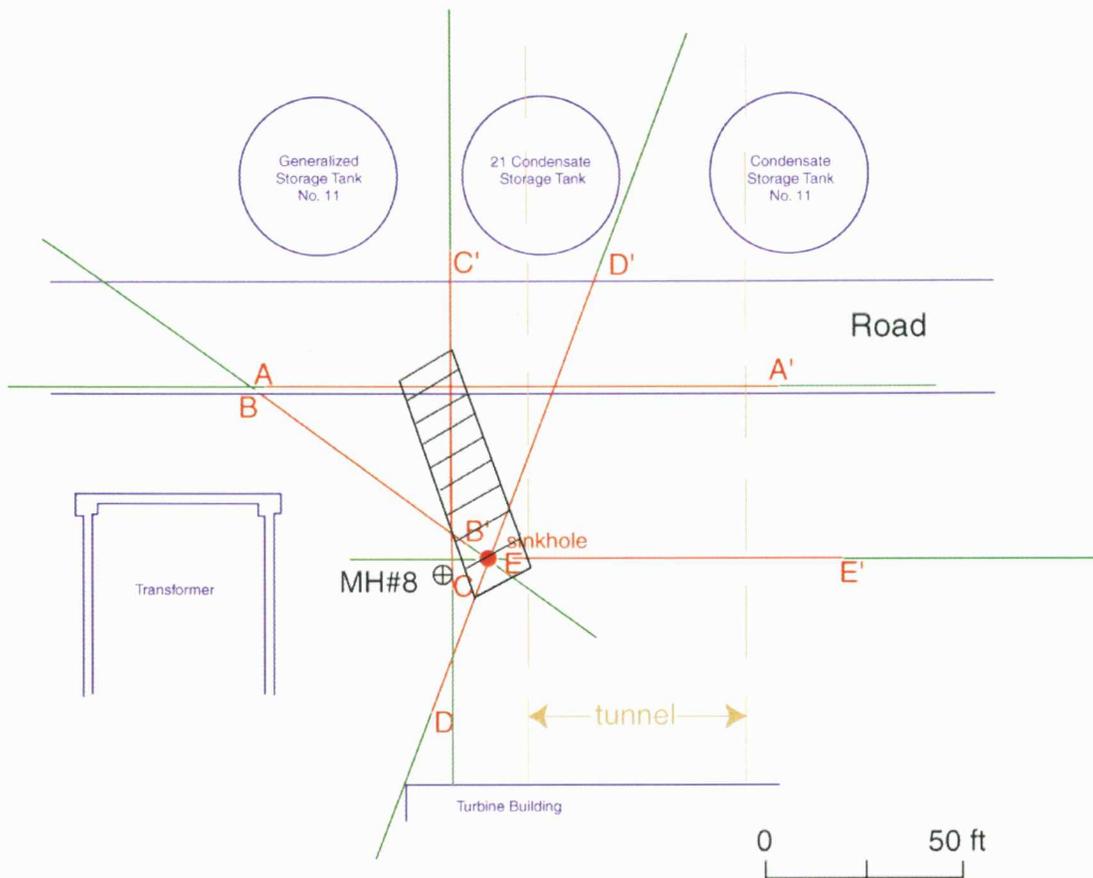


Figure 10. The sinkhole impact area (shaded) interpreted from surface-wave data.

Sinkhole Impact Area

The sinkhole was refilled by soil to prevent possible further damage. Although the refilling reduced the velocity contrast in the sinkhole impact area, we expect that the new soil is much less compacted than surrounding materials so the sinkhole impact area will still show anomalies with a lower S-wave velocity. An imaging point of the surface-wave method is located in the center of the geophone spread of each shot (Miller and Xia, 1999). Because of the limited access to the area surrounding the sinkhole, only line D-D' contains imaging points that traverse the sinkhole. The sinkhole was located at station 155 of line D-D'. A striking relative lower S-wave velocity field from stations 150 to 170 below 40 ft of depth is shown in Figure 9. Based on its low S-wave velocity, it is likely that the area that possesses lower S-wave velocity along line D-D' was impacted by the sinkhole. It is also interesting to see that the sinkhole was located at station 155 instead of 160, the center of the impacted area. This might be due to the smear effect of S-wave velocities at deeper depths. The water tunnel at approximately station 80 on line D-D' showed a higher S-wave velocity. This tunnel might cause a slight shift in the interpreted impacted area on the S-wave velocity field.

Line A-A' (Figure 6) possesses an S-wave velocity field around station 110, similar to that shown on line D-D' around the sinkhole. A chimney-shape S-wave velocity field from

stations 100 to 120 below 40 ft of depth indicate a possible area impacted by the sinkhole. We suggest that the sinkhole impact area extends northward at least to line A-A'. The north-south line C-C' seems to support this suggestion.

The S-wave velocity field of line C-C' (Figure 8) presents an anomalous low that covers the entire line at a depth of 38 ft to 50 ft. This anomaly reveals a possible sinkhole impact area along line C-C'. These results indicate that the chimney-shape anomaly shown on line A-A' was most likely related to the sinkhole.

With a limited accessibility and a limited resolution of the S-wave velocity field, we delineate the sinkhole impact area in Figure 10. Based on the amount of soil required to fill the sinkhole (700 tons = $700 \times 2000 \text{ lb} = 700 \times 900 \text{ kg}$), the volume of the void associated with the sinkhole is estimated to be around 394 m^3 (assuming a density of soil $1,600 \text{ kg/m}^3$). The volume of the void will be larger than 390 m^3 if the soil density is less than $1,600 \text{ kg/m}^3$. The low S-wave velocity zone at depths of 38 to 50 ft along line C-C' (Figure 8) suggest the vertical impact dimension is around 12 ft (3.6 m). We also see the impact length along both lines A-A' and D-D' are around 20 ft (6 m). Thus, a northwest-southeast distance for a interpreted sinkhole impact area is approximately 62 ft (18.5 m), assuming the impact area is approximately a rectangle.

The S-wave velocity field of line E-E' (Figure 7) shows a dramatic S-wave velocity low at station 144 and a depth of 20 ft. This suggests that the area may also be affected by the sinkhole or other voids.

Signatures from Tunnel

Three lines—A-A', D-D', and E-E'—cross the 60-ft wide water tunnel. It was interesting to see how the S-wave velocity field reveals the water tunnel. There was an obvious S-wave velocity increase at depths of around 45 ft at station 145 on line A-A' (Figure 6). The contour line at 4,500 ft/s is associated with the velocity increase. The horizontal location (station 145) and depth (45 ft) of the S-wave velocity increase correspond to the spatial location of the west side of the tunnel. We see the same results on line E-E' (Figure 7); the contour line at 4,250 ft/s is associated with the S-wave velocity increase and suggests the east side of the tunnel. That the tunnel was revealed by the contour lines with different velocity values provides an indication of the resolution of the surface-wave method in this environmental setting. Due to the noise level, the difference of 250 ft/s in S-wave velocity maps did not surprise us.

It is interesting to note that the S-wave velocity field reveals only one side of the tunnel. Based on the data-acquisition geometry, the source was on the west side of the geophone spread for line A-A' and on the east side of the geophone spread for line E-E'. The data-acquisition geometry and S-wave velocity field suggest that surface waves change their behavior significantly in response to S-wave velocities of materials when they propagate from a normal zone and then cross to an anomalous zone. Surface-wave fields, however, may not be sensitive enough to reveal the edge of the anomalous zone and the normal zone when surface waves cross over the anomalous zone and enter the normal zone. This might be due to low seismic energy at far offset geophones so that surface waves are not able to detect the other side of tunnel. We will verify this phenomenon in future studies.

We see little indication of the existence of the tunnel from line D-D' (Figure 9). The sources were on the north side of the geophone spread for this line. The surface is covered by gravel on the north side of the road. It was difficult to generate seismic energy and also difficult for seismic energy to propagate on this portion of the line. In addition to the high-noise environment, this may explain why the S-wave velocity field did not reveal the tunnel.

Conclusions

The surface-wave survey was conducted with limited resources: a 24-channel seismograph with a hammer and an aluminum plate as a source. Although the surface-wave survey at Calvert Cliffs Nuclear Power Plant was conducted at a noise level 50-100 times higher than a normal environment for a shallow seismic survey, the S-wave velocity field calculated from surface-wave data delineated a possible sinkhole impact area. Depending on the source-geophone geometry, the S-wave velocity field revealed one side of the tunnel. Future study should examine why only one side of the tunnel was revealed on the S-wave velocity maps.

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