

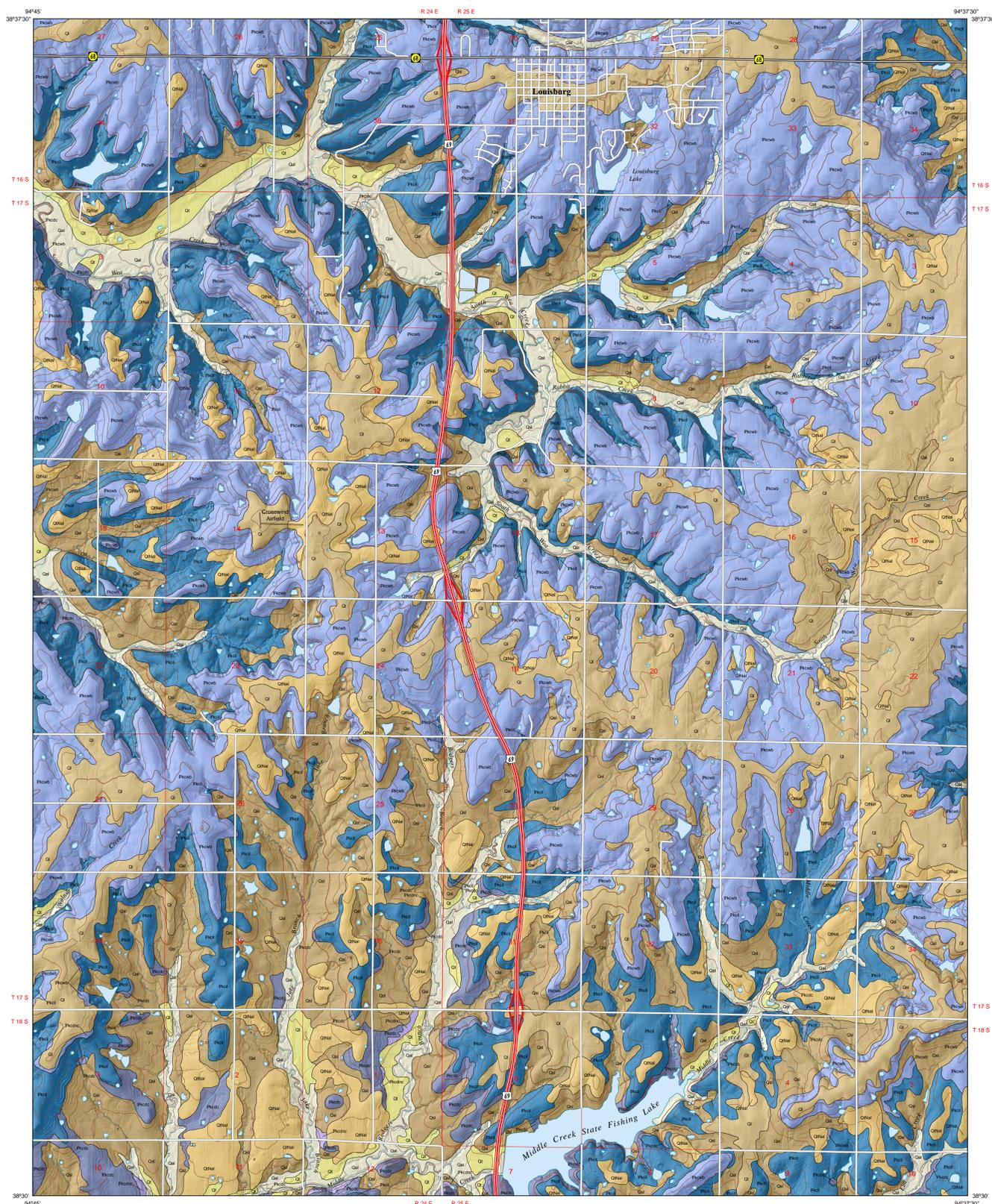
# PRELIMINARY SURFICIAL GEOLOGY OF THE LOUISBURG QUADRANGLE, MIAMI COUNTY, KANSAS

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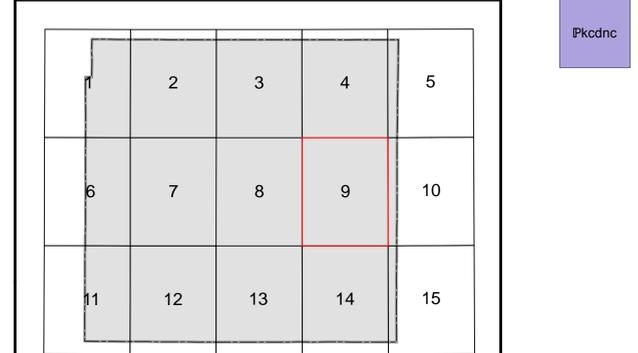
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- GEOLOGIC UNITS**  
CENOZOIC  
Quaternary System
- Qal** — Undifferentiated alluvium — Undifferentiated alluvium that typically is less than 4,000 years old comprises valley fills beneath the modern floodplain of the Marais des Cygnes River and its tributaries. This alluvium consists of clay, silt, sand, and gravel, with the coarsest sediments comprising the lower 3-6 ft (0.9-1.8 m) of the valley fills. The fine-grained floodplain facies are mostly dark gray, gray, dark grayish brown, and grayish brown silty clay, and silty clay loam. These deposits can reach thicknesses greater than 50 ft (15 m) in the Marais des Cygnes River valley.
  - Qt** — Alluvial terraces — Alluvium that typically is 4,000 to 11,000 years old comprises valley fills beneath a low terrace (T-1) of the Marais des Cygnes River and its tributaries. This alluvium consists of clay, silt, sand, and gravel, with the coarsest sediments comprising the lower 3-6 ft (0.9-1.8 m) of the terrace fills. The fine-grained alluvium is mostly brown, dark yellowish brown, and yellowish brown silty clay loam. These deposits can reach thicknesses greater than 70 ft (21 m) in the Marais des Cygnes River valley.
  - QtNal** — Quaternary and Neogene alluvium — Clayey alluvium often interbedded with cherty gravel is common beneath high terraces and on hilltops in Miami County. The numerical age of these alluvial deposits is unknown. However, based on their position in the landscape, the deposits on the hilltops probably aggraded during the Neogene (Aber, 1998), and it is likely that the high-terrace fills date at least to the middle Pleistocene. The lower 3-6 ft (0.9-1.8 m) of the terrace fills and hilltop deposits consist of cherty gravel. The overlying fine-grained alluvium mostly consists of dark gray, gray, dark grayish brown, and pale brown silty clay. The distribution of fine-grained alluvium interbedded with cherty alluvial gravels is indicated by the Kenoma soil series, as shown on the Soil Survey of Linn and Miami Counties, Kansas (Penner, 1981). Redoximorphic features, including strong brown and reddish brown mottles and iron and manganese oxide concretions, are common in the upper 3-10 ft (0.9-3 m) of the fine-grained alluvium. Alluvial deposits comprising the fills of high terraces are typically 20-30 ft (6-9 m) thick, but the alluvial deposits on hilltops generally are less than 15 ft (4.5 m) thick. A veneer of loess that is less than 3 ft (0.9 m) thick often caps the high-terrace and hilltop alluvial deposits.
  - Qsl** — Colluvial apron and alluvial fan deposits — Deposits of clay, silt, sand, and gravel occur on footslopes and toeslopes in valley landscapes. These deposits mostly accumulated during the early and middle Holocene, although some may date back to the terminal Pleistocene, and they comprise both colluvial aprons and alluvial fans. The colluvial aprons consist of massive deposits of poorly sorted sediments that include many angular, bedrock-derived pebbles and cobbles. The alluvial fans formed where small, intermittent streams enter the Marais des Cygnes River valley and the valleys of its major tributaries. The fan deposits are stratified and consist of well-sorted alluvium dominated by brown, dark yellowish brown, and yellowish brown silty clay, silty clay loam, and silt loam. Thin lenses of gravel are common, and buried soils often occur at the top of upward-fining sequences.
  - Ql** — Loess — Deposits of grayish brown and reddish-brown loess cover hilltops and high alluvial terraces in portions of Miami County. Deposits of grayish brown silt loam comprise the Peoria Loess, which dates between ca. 20,000 and 12,000 B.P. The distribution of the Peoria Loess is indicated by the Grundy soil series. In some areas erosion has stripped off the Peoria Loess and exposed reddish brown silty sediment comprising the Loveland Loess. The distribution of the Loveland Loess is indicated by the Welda soil series. The Loveland Loess accumulated between ca. 160,000 and 130,000 B.P. Deposits of loess greater than 10 ft (3 m) thick were mapped.

- EXPLANATION**
- |                                     |   |
|-------------------------------------|---|
| <b>Boundaries and Locations</b>     | <b>Hydrology and Topography</b>         |
| — Township/Range line               | — Perennial stream                      |
| — Section line                      | — Intermittent stream                   |
| <b>Transportation</b>               | — Water body                            |
| — U.S. highway                      | — Water body - manmade shoreline        |
| — State highway (100-foot interval) | — Elevation contour (100-foot interval) |
| — Local road                        | — Elevation contour (20-foot interval)  |
| — Airport runway                    | — Depression contour (20-foot interval) |
| <b>Geologic Unit Boundaries</b>     | — Depression contour (30-foot interval) |
| — Observed contact                  |   |



- MIAMI COUNTY QUADRANGLES**
- |                   |                    |
|-------------------|--------------------|
| 1 Wellsville      | 9 <b>Louisburg</b> |
| 2 Antioch         | 10 Freeman MO-KS   |
| 3 Spring Hill     | 11 Lane            |
| 4 Bucyrus         | 12 Osawatometie    |
| 5 West Line MO-KS | 13 Fontana         |
| 6 Rantoul         | 14 New Lancaster   |
| 7 Paola West      | 15 Drexel MO-KS    |
| 8 Paola East      |                    |

Elevation contours are presented for general reference. Used in the U.S. Geological Survey's current US Topo 1:24,000-scale topographic map series, they were generated from hydrographically-improved 1/3 arc-second National Elevation Dataset (NED) data, and smoothed during processing for use at 1:24,000 scale. In some places the contours may be more generalized than the base data used for compilation of geologic outcrop patterns. Outcrop patterns on the map will typically reflect topographic variation more accurately than the associated contour lines. Repeated fluctuation of an outcrop line across a contour line should be interpreted as an indication that the mapped rock unit is maintaining a relatively constant elevation along a generalized contour.

Geology was mapped in the field using a USGS 7.5-minute 1:24,000-scale topographic map.

Roads and highways are shown on the base map as represented by data from the Kansas Department of Transportation (KDOT), U.S. Census Bureau, and other sources. U.S. Department of Agriculture - Farm Services Agency (USDA-FSA) National Agriculture Imagery Program (NAIP) imagery also was used to check road locations.

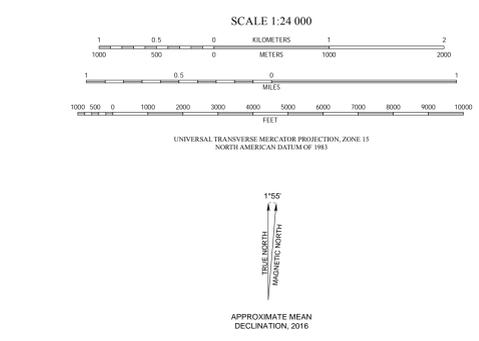
Shaded relief is based on 1-meter hydroflattened bare-earth DEMs from the State of Kansas LIDAR Database. The DEM images, in ERDAS IMAGINE format, were mosaicked into a single output DEM, downsampled to 5-meter resolution, and reprojected to decimal degrees. The output DEM was then converted to a hillshade, a multidirectional shaded-relief image using angles of illumination from 0°, 225°, 270°, and 315° azimuths, each 45° above the horizon, with a 4x vertical exaggeration.

This geologic map was funded in part by the USGS National Cooperative Geologic Mapping Program, award number G15AC00225 (FY2015).

This map was produced using the ArcGIS system developed by Esri (Environmental Systems Research Institute, Inc.).

This map is a preliminary product and has had less scientific and cartographic review than the Kansas Geological Survey's M-series geologic maps. KGS does not guarantee this map to be free from errors or inaccuracies and disclaims any responsibility or liability for interpretations made from the map or decisions based thereon.

**SUGGESTED REFERENCE TO THE MAP**  
Mandel, R. D., and Newell, K. D., 2016, Preliminary surficial geology of the Louisburg quadrangle, Miami County, Kansas: Kansas Geological Survey, Open-File Report 2016-12, scale 1:24,000, unpublished.



**Pkcbw** — **Wyandotte Limestone/Bonner Springs Shale** — The uppermost part of the Kansas City Group is composed of the Wyandotte Limestone and the overlying Bonner Springs Shale. The Wyandotte Limestone is composed of five alternating shale and limestone members. From the base upward, these five units are the Frisbie limestone, Quindaro shale, Argentine limestone, Island Creek shale, and Farley limestone. The Wyandotte is a cuesta-forming formation, ranging in thickness from about 10 to 80 ft (3 to 24 m). The shale members are present only locally in Miami County. The absence of the shale members and the similarity in the lithology of the two upper limestone members makes identification of the units difficult. The **Frisbie Limestone Member** (2.5-3.6 ft; 0.8-1.1 m) is a light olive-gray to light brownish-gray, fine-grained, massive limestone. Locally, thin shales (<0.5 ft; 0.15 m thick) can be present in the Frisbie. Crinoid stems and small productid brachiopods are the most common fossils and locally *Osagia*-like forms (coated grains) may be present. The **Quindaro Shale Member** (locally absent to 3.3 ft; 1 m) is a dusky-yellow to dark yellowish-orange, sandy shale. In places the lower part of this unit is composed of very dark gray shale. The Quindaro is very fossiliferous with *Heterocoelia* sp., *Dielsma* sp., and *Phricodothis* sp. being the most common forms. Crinoid fragments and bryozoans are also abundant. The **Argentine Limestone Member** (~35 ft; 11 m) is probably the most persistent member of the Wyandotte. It is lithologically similar to the Farley limestone which is normally the uppermost limestone member of the Wyandotte. Locally the Farley limestone can be missing and the Argentine limestone is the uppermost member. The Argentine is a light olive-gray to grayish-orange, medium-grained, thin-bedded, locally cherty limestone, which weathers into thin fragments. The Argentine has a varied fauna with the brachiopods *Composita* sp., *Echinaria* sp., *Antiquatonia* sp., and *Phricodothis* sp. being most common. *Entolites* sp. is abundant west and south of Paola. The faunalid *Trilicites* sp. is common locally. The **Island Creek Shale Member** (~1.5 ft; 0.5 m) is not well exposed. It is a grayish-orange clayey shale containing abundant gastropods. The **Farley Limestone Member** (locally absent to ~15 ft; 4.6 m), where present, is the uppermost member of the Wyandotte Limestone. It is a light olive-gray to pinkish-gray coarse-grained, wavy, thin- to thick-bedded limestone. Locally the lower part of this unit can be oolitic. The Farley limestone contains much the same fauna as the Argentine limestone. The **Bonner Springs Shale** (0.9-3.15 ft; 0.3-0.96 m) is the uppermost formation of the Kansas City Group. The lower part is composed of pale olive- to light-gray sandy shale which in places is a thin-bedded micaceous siltstone. The upper part is an olive-gray to yellowish-brown clayey shale in the upper part. Varying lithologies can be present near the middle of the unit - unfossiliferous grayish-red clayey shale or medium-gray clayey shale containing carbonaceous streaks, plant roots, and leaf impressions, or medium-grained, medium-bedded, calcareous sandstone is found in this interval. In areas where the Bonner Springs is primarily clayey shale, scattered limestone nodules are present, and a thin, argillaceous limestone bed (~1 ft; 0.3 m) is present 2.0 ft (0.6 m) below the top of the formation. The limestone bed is very fossiliferous and contains abundant pelecypods, algae, brachiopods, gastropods, and bryozoan fragments.

**Pkcl** — **Iola Limestone/Lane Shale** — The **Iola Limestone** (10-12 ft; 3-3.7 m) is composed of two limestone members and an intervening shale member - the lower limestone is the Paola limestone, the middle shale is the Muncie Creek shale, and the upper limestone is the Raytown limestone. The **Paola Limestone Member** (absent to ~2.5 ft; 0.8 m) is composed of a massive bed of dark-gray to brownish-gray, fine-grained, dense limestone. The contact with the underlying shale is fairly smooth, but the upper surface of the unit is very irregular. Locally, iron-stained "worm tubes" extend downward 1 to 3 inches (2.5-7.6 cm) from the upper surface of the Paola. Crinoid stems and small productid brachiopods are the most abundant fossils. Coated grains and bryozoans are common. The **Muncie Creek Shale Member** (0.5 ft; 0.15 m) is a bluish-gray to dusky-yellow sandy shale that locally has a dark-gray, carbonaceous, fissile shale that contains small (0.1-0.2 inches; 1-2 cm) phosphatic nodules. The **Raytown Limestone Member** (5-24 ft; 1.5-7 m) is a light olive-gray to light-gray medium- to coarse-grained limestone with numerous silt shale partings and abundant vugs lined with crystalline calcite. The Raytown is medium bedded in the northern part of Miami County but becomes progressively thin bedded toward the southern part. Abundant large productid brachiopods occur in the Raytown, with *Echinaria* sp. and *Lingulodictya* sp. being the most common. The **Lane Shale** (5-8 m; increasing in thickness westward to 80 ft; 24 m) is lithologically variable. In the western half of the county is an olive-gray to dusky-yellow silty to sandy shale and thin-bedded siltstone. In the eastern half it is an olive-gray to light-gray clayey shale. Thin carbonaceous streaks 1/16- to 1/4-inch (1.6- to 6.4 mm) thick are found locally in the Lane. In east-central Miami County, a thin light-gray, clayey shale containing laminae of reddish-brown calcareous silt occurs just below the overlying Wyandotte Limestone. The Lane is unfossiliferous except for sparse plant remains, and locally, small brachiopods and crinoids are present in the upper few feet of the formation.

**Pkcdc** — **Drum Limestone/Chanute Shale** — The **Drum Limestone** (1.7-8 ft; 0.5-2.4 m) consists of a single massive bed of yellowish-gray to reddish-brown fine- to medium-grained cross-bedded limestone that is uniform in lithology and weathers into thin slabs. The most characteristic feature of this unit is abundant, small, white crinoid segments scattered throughout the limestone. Other fossils are *Neospirifer* sp., *Margarifera* sp., and *Composita* sp., and locally the small sponge *Heliospongia* sp. The **Chanute Shale** (8-38 ft; 2.4-12 m) varies greatly in thickness and lithology over short lateral distances. It is yellowish-brown to greenish-gray sandy to clayey shale, which locally contains sandstone in the lower and middle parts and commonly in the upper part. The Chanute Shale is a relatively unfossiliferous unit except for plant impressions found locally in the sandy layers. The Chanute Shale contains the Thayer coal (0-0.5 ft; 0.15 m), 11 to 15 feet (3.4-4.6 m) above the base of the formation. The Thayer occurs at differing stratigraphic positions in respect to the upper formational boundary, but is absent where the Chanute is thinner than 11 ft (3.4 m).

**Pkcdnc** — **Dennis Limestone/Cherryvale Shale** — The **Dennis Limestone** (~32 ft; 10 m) is resistant, forming prominent scarps. It is composed of three members, which, in ascending order, are the Canville limestone, Stark shale, and Winterset limestone. The **Canville Limestone Member** is absent in Miami County, but may be represented locally by a thin, impure, nodular limestone. The **Stark Shale Member** (0.9-6 ft; 0.3-1.8 m) is represented by a black, fissile to very thin-bedded, carbonaceous shale in the lower part. The upper part is medium-olive to greenish-gray sandy shale and is about twice the thickness of the underlying black shale. The Stark is fossiliferous in the upper, sandy part, with chonetid brachiopods being the most common fossil. The lower part is unfossiliferous except for rare articulate brachiopods. The **Winterset Limestone Member** (~29 ft; 9 m), the uppermost member of the Dennis Limestone, is light gray to olive gray, medium grained, medium bedded, dense in the lower part and more fine grained in the middle part. The lower and middle parts of the member characteristically contain shaly partings and abundant chert. The upper part is usually separated from the lower bed by a thin bed of calcareous gray shale (~1.5 ft; 0.5 m). The upper part of the Winterset is dense, fine-grained, massive limestone. It contains an abundance of dark-gray chert. Locally, 1 to 3 ft (0.3-0.9 m) of oolitic limestone can be present near the top. The upper part of the Winterset is very fossiliferous and is characterized by numerous large productid brachiopods and locally by abundant gastropods. *Trilicites* sp. is also very abundant in the upper part and is usually found in chert. The **Cherryvale Shale** (~60 ft; 18 m) comprises a shale-dominated unit between the top of the Dennis Limestone and the base of the Drum Limestone. It includes five members, in ascending order: the Fontana shale, the Block limestone, the Wea shale, the Westerville limestone, and the Quivira shale. The **Fontana Shale Member** (~15 ft; 4.6 m) is greenish-gray to olive-gray sandy shale. Locally a thin, nodular limestone occurs in the lower part. It is relatively unfossiliferous except for sparse chonetid brachiopods in the lower part. The **Block Limestone Member** (~4 ft; 1.2 m), the lowermost limestone bed in the Cherryvale Shale, is a bluish-gray to olive-gray medium-grained, thin-bedded fossiliferous limestone with numerous thin, fossiliferous shale partings. *Trilicites* sp. is the most characteristic fossil, but *Margarifera* sp. and *Syringopora* sp. also are abundant. The Wea Shale Member (~18 ft; 5.5 m) is characteristically a light gray sandy shale that contains sandstone beds locally near the base. A thin bed of maroon shale may occur near the top. Locally, pyritiferous nodules and limonite concretions occur in the Wea, but plant impressions and carbonaceous partings can be found at its top. In places a very thin coal occurs near the top of the member. The **Westerville Limestone Member** (absent to ~2 ft; 0.6 m) is poorly exposed, but where present, is composed of conglomeratic limestone containing abundant limonite nodules and quartz sand. It can lithologically transition to yellowish-gray marly limestone. The **Quivira Shale** (~4 ft; 1.2 m) has a lower part (0.5-1 ft; 0.15-0.3 m) that is composed of black, carbonaceous, fissile shale or maroon clayey shale. The upper part consists of olive-gray sandy shale. The Quivira is unfossiliferous, except for sparse inarticulate brachiopods and conodonts in the black fissile shale.

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